

## **Compressibility and Permeability Characteristics of Selected Coastal Soils of Bangladesh**

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### **Introduction**

**A** geotechnical engineer is concerned with the compressibility characteristics of soil because he has to know by how much a soil will be compressed under external load and at what rate the compression will take place. On the other hand, permeability governs such important engineering problems as the ground water regime in layered deposits, or near natural and excavated slopes, the consolidation of clay foundations under applied loads and the flow of water through or around engineering structures.

In recent times several coastal regions of Bangladesh have been affected by severe cyclone and storm surges which produced quite significant damage to livestock, agriculture, power system, telecommunication, housing and other physical infrastructure facilities. Flood protection embankments, retaining structures and cyclone shelters are needed to be built in large quantities in these coastal regions. It is essential to understand the mechanical properties and behaviour of these soils, which are in general silty clays of low to high plasticity.

This paper presents the compressibility and permeability characteristics of four coastal soils. Attempt has also been made to evaluate the essential intrinsic properties of these soils.

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## Compressibility and Permeability Characteristics of Clays

Expression for rate of consolidation has been first developed by Terzaghi (1943). Based on Terzaghi's theory an ASTM standard is formulated (1979). Recently, two other one-dimensional consolidation test procedures have been developed which are much faster and provide reasonably good results (Lowe, 1969; Smith and Wahls, 1969). For a given stress increment, the coefficient of consolidation  $c_v$  can be determined from laboratory observations of time versus dial reading. Two graphical procedures are commonly used for this are the logarithmic of time method proposed by Casagrande and Fadum (1940) and the square root of time method proposed by Taylor (1942).

$$c_v = \frac{0.197 H^2}{t_{50}} \quad \text{and} \quad c_v = \frac{0.848 H^2}{t_{90}} \quad (1)$$

where,  $H$  is the average length of drainage path, i.e., half the height of sample at the end of 100% consolidation for a given stress increment. There are two other useful methods, which are proposed by Su (1958) and Sivaram and Swamee (1977).

Several researchers have suggested functions for the relationship between compression index and void ratio for the undisturbed samples of coastal soils of Bangladesh (Serajuddin and Ahmed, 1967; Amin et al., 1987; Kabir et al., 1992). The empirical correlations between  $C_c$  and  $e_0$  for undisturbed samples, as proposed by Serajuddin (1969) and Amin et al. (1987) are given by the following Eqns. (2a) and (2b) respectively:

$$C_c = 0.50(e_0 - 0.50) \quad (2a)$$

$$C_c = 0.33(e_0 - 0.35) \quad (2b)$$

Samarasinghe et al. (1982) defined the void ratio  $e_1$  as the value of void ratio when the effective stress is 1 tsf (95.7 kPa). Samarasinghe et al. (1982) termed  $C_c$  and  $e_1$  as compressibility parameters. The values of  $e_1$ ,  $C_c$ ,  $C_s$  and  $c_v$  obtained from reconstituted soil samples studied by several researchers are summarized in Table 1. Using  $C_c$  and  $e_1$  from Raymond (1966), Samarasinghe et al. (1982) and a number of regional soils of Bangladesh, Siddique (1986) proposed the following relation between  $C_c$  and  $e_1$  for reconstituted samples

$$C_c = 0.39(e_1 - 0.14) \quad (3)$$

**Table 1**  
**Values of LL, PI,  $e_1$ , Compression Index, Swelling Index and Coefficient of Consolidation of some Reconstituted Samples**

Location	LL	PI	$e_1$	$C_c$	$C_s$	$c_v$ ( $10^{-4}$ cm <sup>2</sup> /cm)	
						log t fitting	$\sqrt{t}$ fitting
Don Valley Clay <sup>1</sup>	41	22	0.78	0.26	—	—	—
New Liskeard <sup>1</sup>	69	40	1.14	0.41	—	—	—
Greyish Sandy Clay <sup>2</sup>	27	14	0.57	0.16	—	3 to 7	4 to 9
Gazipur Clay <sup>3</sup>	33	6	0.93	0.29	—	70 to 180	100 to 250
London Clay <sup>4</sup>	69	45	1.16	0.55	0.15	0.8 to 4	1 to 6
Dhaka Clay <sup>5</sup>	40	20	0.84	0.28	—	5 to 10	7 to 14

1: Raymond (1966); 2: Samarasinghe et al. (1982); 3: Siddique (1986);  
 4: Siddique and Clayton (1994); 5: Siddique and Safiullah (1995)

Burland (1990) introduced a new normalizing parameter termed as the void index to aid in correlating the compression characteristics of various clays. Later Kamaluddin (1990) established the same relationship for Dhaka clay. Burland (1990), termed the properties of reconstituted clays as 'intrinsic' properties since they are inherent to the soil and independent of natural state. The term 'intrinsic' has been used to describe the properties of clays which have been reconstituted at a water content of between  $w_L$  and  $1.5w_L$  and then consolidated under one-dimensional condition.

The compressibility and strength characteristics of reconstituted clays were used as a basic frame of reference for interpreting the corresponding characteristics of natural sedimentary clays. The intrinsic properties provide a frame of reference for assessing the in-situ state of a natural clay and the influence of structure on its in-situ properties. One dimensional compression curve for reconstituted natural clay is normalized by assigning fixed values to  $e_{100}^*$  and  $e_{1000}^*$ . The parameters  $e_{100}^*$  and  $e_{1000}^*$  are the intrinsic void ratios corresponding to consolidation pressures,  $\sigma'_v = 100$  kPa and  $\sigma'_v = 1000$  kPa, respectively. The normalized parameter chosen has been defined as the void index,  $I_v$  such that,

$$I_v = \frac{e - e_{100}^*}{e_{100}^* - e_{1000}^*} = \frac{e - e_{100}^*}{C_c} \quad (4)$$

The intrinsic compression index,  $C_c^*$  is defined as the difference in void ratio between  $e_{100}^*$  and  $e_{1000}^*$ .

Following Terzaghi (1925), the parameters  $e_{100}^*$  and  $C_c^*$  are called as the constant of intrinsic compressibility. When  $e = e_{100}^*$ , then  $I_v = 0$  and when  $e = e_{1000}^*$ , then  $I_v = -1$ . The void index may be thought of as a measure of the intrinsic compactness of a sediment. When  $I_v$  is less than zero, the sediment is compact and when  $I_v$  is greater than zero, the sediment is loose. It has been seen that a reasonably unique line is achieved when void index,  $I_v$  versus  $\log v$  is plotted for clays covering a wide range of liquid limits and of pressures. This line has been termed as intrinsic compression line (ICL). Burland has given a cubic equation as follows which may represent the coordinates of the ICL with sufficient accuracy

$$I_v = 2.45 - 1.285x + 0.015x^3 \quad (5)$$

where  $x = \log p'_0$  in kPa

The intrinsic compression line may either be measured directly for a clay or if the values of  $e_{1000}^*$  and  $C_c^*$  are known for the clay, the ICL may be constructed using the above Eqn. (5). If it is required to plot the ICL in terms of  $e$  versus  $\log p'_0$ , then the values of  $e$  corresponding to various values of  $\log p'_0$  may be obtained from the following expression:

$$e = I_v C_c^* + e_{100}^* \quad (6)$$

In the above Eqn. (6), the values of  $I_v$  may be obtained from Eqn. (5).

For coarse grained soils, permeability has been related to grain size distribution. Similar relationship for fine-grained soils is less successful to predict permeability accurately for engineering purposes. In the past, numerous researchers have suggested functions for the relationship between permeability and void ratio (Taylor, 1948; Raymond, 1966; Lambe and Whitman, 1969). Recently, Samarasinghe et al. (1982) suggested a model to predict permeability of normally consolidated reconstituted clays. The relationship is as follows:

$$k = C \left( \frac{e^n}{1+e} \right) \quad (7)$$

Samarasinghe et al. (1982) termed  $n$  and  $C$  as permeability parameters. A plot of  $\log[k(1+e)]$  versus  $\log e$  results in a straight line of which  $n$  is the slope and  $\log C$  is the intercept. Samarasinghe et al. (1982) carried out both direct permeability tests and incremental loading consolidation tests on an artificially sedimented normally consolidated sandy clay to verify the model represented by Eqn. (7). For any method,  $n$  is about the same for each soil.

**Table 2**  
**Values of Permeability Parameters of Different Soils Derived from Consolidation Test**

Soil	Direct Permeability Method		log t Fitting Method		$\sqrt{t}$ Fitting Method	
	n	C ( $10^{-10}$ m/sec)	n	C ( $10^{-10}$ m/sec)	n	C ( $10^{-10}$ m/sec)
Don Valley Clay <sup>1</sup>	4.8	29	4.8	15	—	—
New Liskeard Clay <sup>1</sup>	4.9	4	4.7	3.6	—	—
Greyish Sandy Clay <sup>2</sup>	5.2	77	—	—	—	—
Gazipur Clay <sup>3</sup>	5.1	355-	4.8	186	5	282
Dhaka Clay <sup>4</sup>	4.7	12	4.5	8.5	4.8	10

1: Raymond (1966); 2: Samarasinghe et al. (1982); 3: Siddique (1986);  
 4: Siddique and Safiullah (1995)

However, C varies with the method of tests. Table 2 shows the values of n and C obtained for a number of clays.

Raymond (1966) also conducted hydrostatic consolidation tests and direct permeability tests on the same specimen for three different reconstituted normally consolidated clays of medium to high plasticity. It has been observed that the reconstituted clays investigated by Raymond (1966) behave exactly in conformity with Eqn. (7) as proposed by Samarasinghe et al. (1982).

For both reconstituted and natural clays linear plots of void ratio versus logarithm of permeability have also been suggested by other researchers (Tavenas et al., 1983; Tababa and Wood, 1987; Leroueil et al., 1992; Nagaraj et al., 1993). Tavenas et al. (1983) assessed the applicability of Eqn. (7) for a number of undisturbed natural soft clays. The experimental data obtained in the investigation suggest that the relation,  $\log e - \log [k(1+e)]$  is suitable to represent permeability - void ratio relation in natural clays.

## Soils Used

Disturbed samples were collected from four selected locations of coastal region from shallow depths. Undisturbed block samples were also collected from these locations. Figure 1 shows the location of the studied sites. From borehole records and from inspection of open cuts nearby, it was established that the sub-soil at this locations are more or less uniform in the upper layer. Table 3 shows the index properties and classification of coastal soils used.

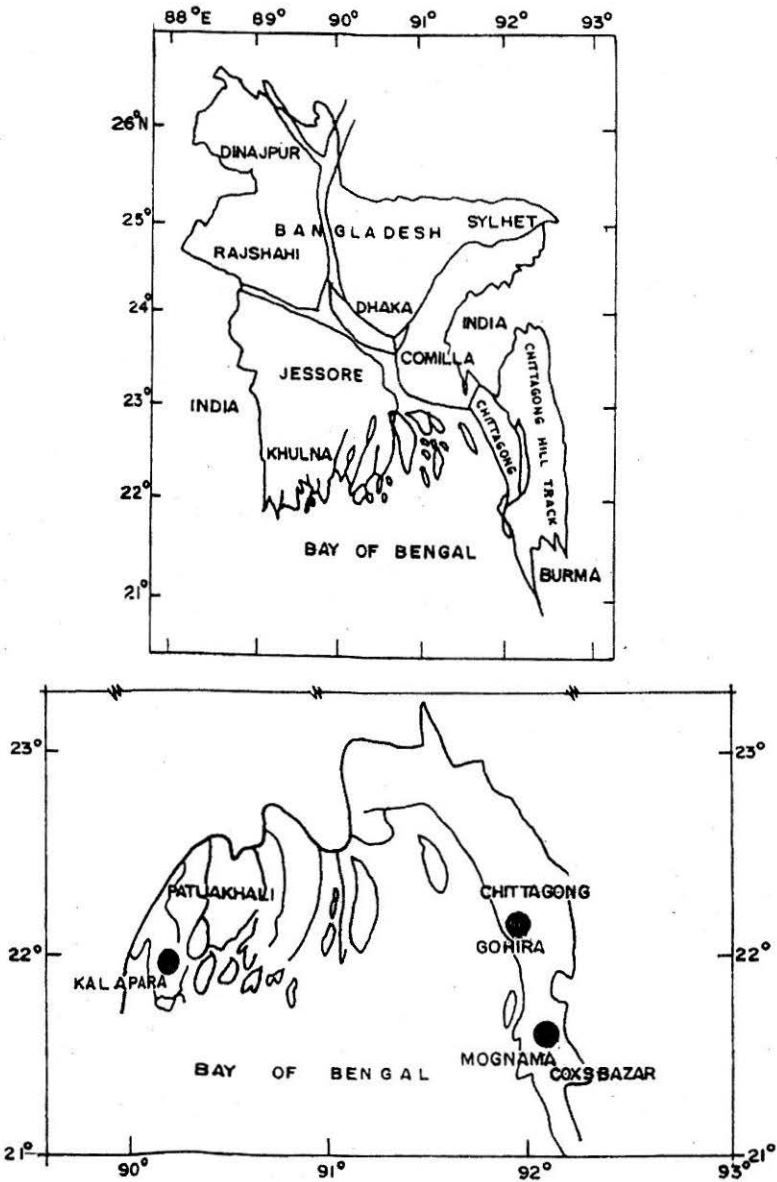


FIGURE 1 : Map of Bangladesh showing the Location of Coastal Sites

**Table 3**  
**Index Properties and Classification of the Coastal Soils Used**

Location	Sample Desig.	$G_s$	LL %	PI %	Clay %	Silt %	Sand %	USCS symbol
Gohira (1.5m) Chittagong	G1	2.79	48	24	41	58	1	CL
Gohira (3.5m) Chittagong	G3	2.71	36	16	38	61	1	CL
Kalapara (GL) Patuakhali	KA	2.65	45	18	30	70	0	ML
Mognama (1m) Cox's Bazar	M	2.73	49	26	39	56	5	CL

USCS: Unified Soil Classification System

## Regional Geology

The coastal region comprising Barisal, Chittagong, Cox's Bazar, Khulna, Noakhali and Patuakhali are underlain by floodplain and meandering deposits laid down by the rivers and their tributaries. The river gradient is flat, and is one of the causes for the reduction of the velocity and deposition of fine eroded particles like silt and clay. The formation is subjected to flooding to a varying depth. The soil profile is generally stratified but their continuity is limited to short distances. Coastal soils can not be classed either as recent or terrace deposits. These are usually estuarine and tidal deposits.

Ground water in the coastal area is strongly influenced by saline water. Tube wells of more than 200 m deep are dug in and around the Chittagong hills and hills near Moheskhal Island to avoid saline water intrusion. Some flowing artesian wells are also observed in these areas. The well depth in other parts of the coastal area of around 300 m is generally much deeper. Some wells near Noakhali are more than 400 m in depth (MCSP, 1992). Most soil studies in the coastal area in the past have dealt with the soil types upto approximately 20 m below the ground surface. The surface layer mainly consists of silt and clay and has a thickness of some 50 m, except at the mouth of the Meghna river where the thickness is reduced to some 10 m. A more detailed examination reveals that the soil texture of the surface layer differs from one area to another in both the horizontal and vertical directions. The grain size, density and consistency also largely differ from one area to another. These differences reflect the sedimentation environment and are caused by frequent changes of the well-developed river and water channel courses. In general, the deposits of the major rivers are coarser than those of the sea currents. A generalized stratigraphic succession of rocks in the area is presented in Table 4.

**Table 4**  
**General Stratigraphic Sequence of Rocks In Coastal Region (Khan, 1991)**

Approximate age	Group	Formation	Lithology
Recent	.....	Beach sand Coastal deposit Alluvium	Sand containing heavy mineral Sand, silt and clay Silt, clay and fine sand
		Unconformity	.....
Mio-Pliocene	Dupi Tila	.....	Coarse ferruginous soft sandstone; clay (sandy and mottled) at intervals; pebbly sandstone and conglomerate beds
		Unconformity	.....
Upper Miocene	Tipam	Girujan Clay	Dark grey and bluish grey clay with mottlings, lower part consists of ferruginous sandstone
		Tipam Sandstone	Coarse and gritty ferruginous sandstone with interlayer of clay; sandstones are massive, cross bedded and yellowish brown in colour; conglomerate in pocket, fossil wood and lignite are present
Middle Miocene	Boka Bil	Upper Unit	Bluish grey shales with conchoidal fractures; lenses of hard calcareous sandstone and shale
		Middle Unit	Alternations of massive and bedded sandstone with thin layers of shales and siltstones; Lenticular and spherical concretions; Conglomerates are occasionally found
		Surma	Lower Unit
Lower Miocene	Bhuban	Unconformity	.....
		Upper Unit	Shales predominate in the lower part with progressively more sand intercalation towards the top: Fine sandstone in the uppermost part
		Middle Unit	Sandstone with subordinate shale Mainly an alternation of shales, sandy shales, and fine grained sandstone
		Lower Unit	Mainly sandy shale, pebbly sandstone and silty shale
		Base not known	.....



## Preparation of Reconstituted Sample

Reconstituted soils are those which are prepared by breaking down natural soils, mixing them as slurry and consolidating them. Reconstituted soil enables a general pattern of behaviour to be established (Jardine, 1985). Reconstituted samples of the four coastal soils were prepared in the laboratory by  $K_0$ -consolidation of a uniform slurry of the clay in a cylindrical consolidation cell of 153 mm diameter and 178 mm height using a consolidation pressure of 100 kPa to get a clay of soft consistency. The slurry had a water content of approximately 1.5 to 2 times the liquid limit of the soil.

## Test Results And Discussion

Consolidation tests were carried out using a stress increment ratio of 1, (i.e., a load ratio of 2) was used. The vertical stresses applied during consolidation were 26.8 kPa, 53.6 kPa, 107.2 kPa, 214.5 kPa, 429 kPa, 858 kPa and 1287 kPa. The samples were also allowed to swell under stresses of 322 kPa and 10.7 kPa. Duration of each loading step was twenty four hours. During all these tests drainage was permitted from top and bottom of the sample. These tests were carried out in accordance with the procedure specified in ASTM (1979) Standards. For each loading step deformation was recorded by a dial gauge at specified intervals of time.

Time – deformation curves are plotted for each pressure increment and from these plots times corresponding to 50% consolidation i.e.,  $t_{50}$  and 90% consolidation, i.e.,  $t_{90}$  were determined using Casagrande's Curve Fitting Method (Casagrande and Fadum, 1940) and Taylor's Curve Fitting Method (Taylor, 1942) respectively. Coefficients of consolidation,  $c_v$ , were calculated for each stress increment using the expression shown in Eqn. (1).

Coefficient of permeability of the samples were determined indirectly from one-dimensional consolidation tests. The coefficient of permeability,  $k$  was computed using the following expression (Taylor, 1948):

$$k = c_v m_v \gamma_w \quad (8)$$

where  $m_v$  = coefficient of volume compressibility, and  
 $\gamma_w$  = unit weight of water.

$m_v$  can be estimated from the following equation:

$$m_v = \frac{C_c}{2303(1+e)\sigma'} \quad (9)$$

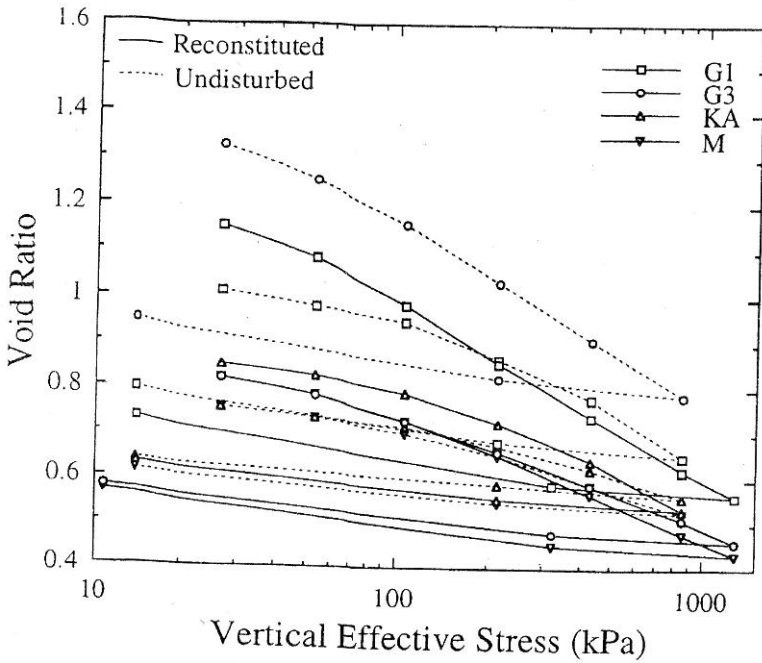


FIGURE 2 : Comparison of Void Ratio Vs. Effective Overburden Pressure Curves between Reconstituted and Undisturbed Samples of Four Coastal Sites

In the following sections compressibility and permeability characteristics of reconstituted and undisturbed samples of four coastal soils obtained from one dimensional consolidation test results are presented and discussed.

### *Compressibility And Expansibility Characteristics*

The compressibility characteristics of reconstituted and undisturbed samples of the four coastal soils undergoing incremental loading in an oedometer are presented in Fig. 2 in which void ratios (at the end of each loading and unloading stages) versus  $\log \sigma_{vo}'$ , (logarithmic of vertical consolidation stress) curves are shown. The values of compression index ( $C_c$ ) and swelling index ( $C_s$ ) have been determined from the loading and unloading portion respectively of the  $e - \log \sigma_{vo}'$  curves. The values of  $C_c$  and  $C_s$  for the four coastal soils are listed in Table 5. It can be seen from Table 5 that the values of  $C_c$  and  $C_s$  for reconstituted samples varied from 0.29 to 0.40 and 0.028 to 0.046, and for undisturbed samples varied from 0.22 to 0.42 and 0.037 to 0.062 respectively. It can also be seen from Table 5 that for samples of Gohira which has high natural water content the  $C_c$  values of the undisturbed samples are higher than those of reconstituted samples. For the samples from Kalapara and Mognama which have comparatively low natural

Table 5  
Values of  $w_n$ ,  $w_1$ ,  $e_1$ , Compression Index, Swelling Index of Four Coastal Soils

Site	Undisturbed Sample				Reconstituted Sample			
	$w_n$ (%)	$e_1$	$C_c$	$C_s$	$w_1$ (%)	$e_1$	$C_c$	$C_s$
G 1	38	0.95	0.42	0.054	37	1.02	0.40	0.046
G3	52	1.17	0.42	0.062	27	0.73	0.29	0.028
KA	28	0.72	0.22	0.043	30	0.80	0.37	0.041
M	26	0.71	0.22	0.037	27	0.73	0.29	0.031

water content, however, the  $C_c$  values of the undisturbed samples are lower than those of reconstituted samples. It is also evident from Table 5 that  $C_s$  of undisturbed samples of Gohira are considerably higher than those of reconstituted samples while the  $C_s$  values of undisturbed samples of Kalapara and Mognama are slightly higher than that of reconstituted samples. Farooq (1995) investigated the compressibility properties of three reconstituted soils from Chittagong belt. The values of  $C_c$  and  $C_s$  reported by Farooq (1995) varied from 0.26 to 0.30 and 0.02 to 0.03 respectively.

Figure 3 shows the compressibility and expansibility characteristics of reconstituted and undisturbed samples of a typical coastal soil (G1: Gohira 1.5m) undergoing incremental loading in an oedometer. In this figure, coefficient of volume compressibility,  $m_v$  and coefficient of volume increase,  $m_s$  have been plotted against  $\log \sigma_{vo}'$ . It can be seen from Fig. 3 that in case of reconstituted sample, for loading up to consolidation stress of 300 kPa,  $m_v$  changes significantly. However, above 300 kPa,  $m_v$  decreases slightly with increasing consolidation stress. During unloading from 1287 kPa to 13.5 kPa, the values of  $m_s$  increased. For reconstituted London Clay (LL = 69, PI = 45), Siddique and Clayton (1994) found that  $m_v$ -values reduced sharply beyond preconsolidation pressure and  $m_s$ -values increased during unloading. Farooq (1995) found  $m_v$  values between  $1.13 \times 10^{-4}$  to  $6.1 \times 10^{-4}$  cm<sup>2</sup>/gm for three reconstituted samples of coastal soils from Chittagong. Figure 3 also shows the relation between coefficient of volume compressibility and volume expansibility with effective vertical stress for undisturbed sample. It is evident from Fig. 3 that the values of  $m_v$  and  $m_s$  of reconstituted sample are comparatively higher than those for the undisturbed sample.

Figure 4 shows typical plottings of coefficient of consolidation,  $c_v$  versus vertical effective consolidation stress for reconstituted and undisturbed samples of Gohira. It is also evident from Fig. 4 that for the reconstituted samples  $c_v$  value generally increased with increasing levels of consolidation stress. Increase in  $c_v$  with increasing levels of vertical effective consolidation stress has been reported by Samarasinghe et al. (1982) and Siddique (1986) for

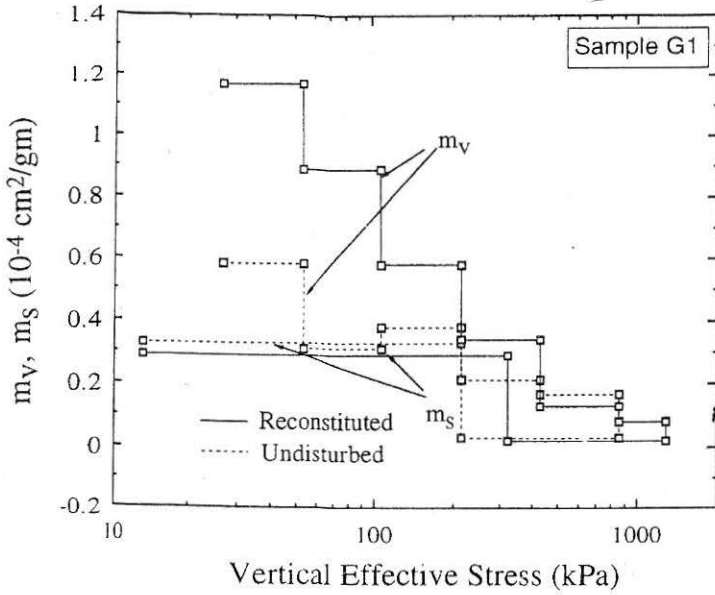


FIGURE 3 : Coefficient of Volume Compressibility and Volume Expansibility Vs. log of Vertical Effective Stress Plots for a Typical Coastal Soil

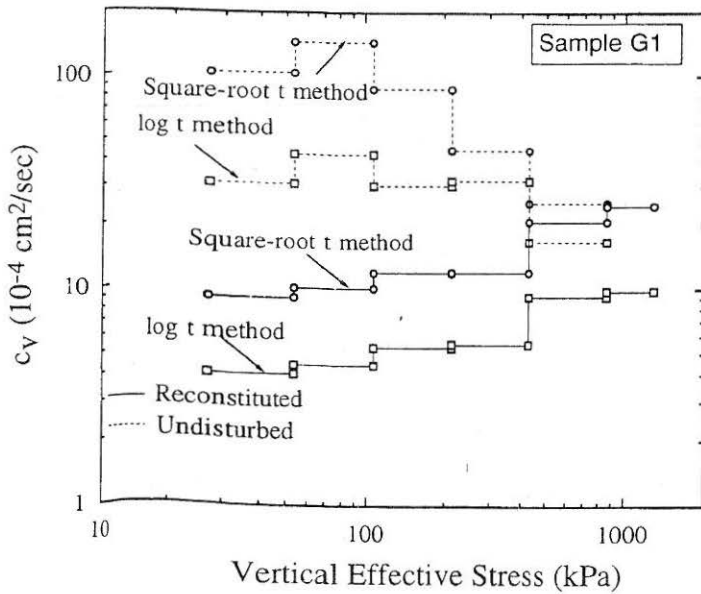


FIGURE 4 : Coefficient of Consolidation Vs. log of Vertical Effective Stress Plots for Reconstituted and Undisturbed Samples of a Typical Coastal Soil

reconstituted normally consolidated clays. Figure 4 shows that for both reconstituted and undisturbed samples, the values of  $c_v$  obtained using  $\log t$  method are less than those obtained using  $\sqrt{t}$  method. Figure 4 also shows that for the undisturbed samples, the values of  $c_v$ , in general, reduced with the increase in consolidation stress. The values of  $c_v$  for the undisturbed samples have been found to be significantly higher than those for reconstituted samples. The values of  $c_v$  for the four reconstituted and undisturbed samples of the coastal soils varied from  $3.3 \times 10^{-4}$  to  $30.2 \times 10^{-4}$   $\text{cm}^2/\text{sec}$  and  $10.1 \times 10^{-4}$  to  $224.4 \times 10^{-4}$   $\text{cm}^2/\text{sec}$  respectively. Serajuddin (1969) reported  $c_v$  values between  $0.2 \times 10^{-4}$  to  $200 \times 10^{-4}$   $\text{cm}^2/\text{sec}$  for large number of undisturbed samples of coastal soils while Farooq (1995) reported  $c_v$  values between  $20 \times 10^{-4}$  to  $57 \times 10^{-4}$   $\text{cm}^2/\text{sec}$  for three reconstituted samples of Chittagong coastal region.

Attempts have also been made to relate compressibility parameters  $C_c$  and  $e_1$  for the reconstituted samples. Based on regression analysis of these data the following correlation between  $C_c$  and  $e_1$  has been derived:

$$C_c = 0.384(e_1 + 0.06) \quad (10)$$

The above relationship is presented in Fig. 5. In Fig. 5 the relation between  $C_c$  and  $e_1$  as proposed by Siddique (1986) and given in Eqn. (3) is

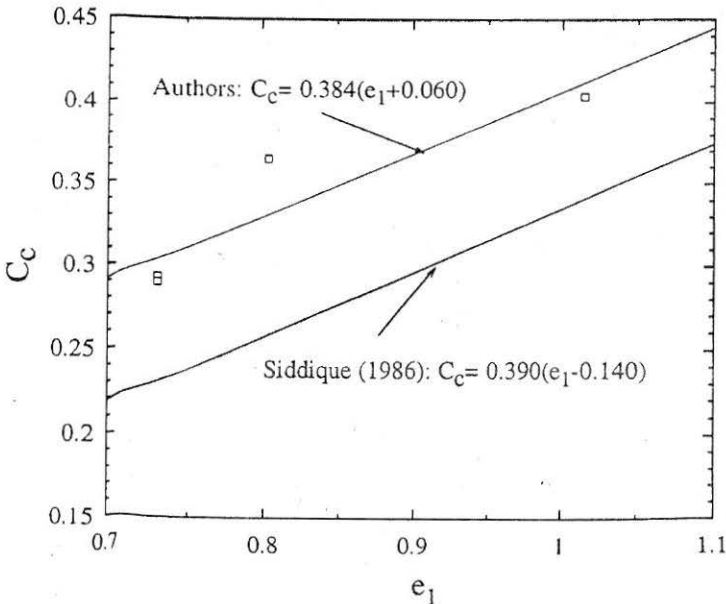
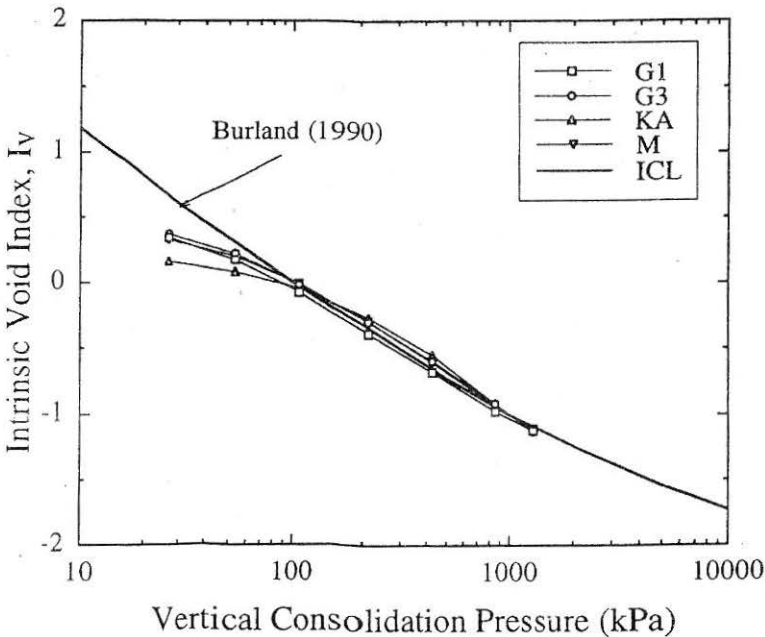


FIGURE 5 : Plot of  $C_c$  Vs.  $e_1$  for Reconstituted Samples

also shown for comparison. It can be seen from Fig. 5 that the line representing relationship between  $C_c$  and  $e_1$  obtained for the soils studied is parallel to that proposed by Siddique (1986) for a number of reconstituted samples and lies above the line proposed by Siddique (1986).

***Void Index and Intrinsic Compression Lines***

The compression curves shown in Fig. 2 represent the intrinsic compression curves for the four locations of coastal region, since the samples were reconstituted at water content of 1.5 times the liquid limit. It is useful to normalize these laboratory compression curves with respect to void ratio. Using Burland's concept (1990), the intrinsic compression curves for the coastal regions shown in Fig. 2 have been normalized by plotting void index,  $I_v$ , versus  $\log \sigma'_v$  and which are shown in Fig. 6. The intrinsic void ratio  $e_{100}^*$ , intrinsic compressibility,  $C_c^*$ , void ratio at the liquid limit,  $e_L$ , the equation of void index for the four locations of the coastal region are shown in Table 6. Using the equations given for the void index in Table 6, intrinsic compression lines for the four coastal soils are shown in Fig. 6. It can be seen from Fig. 6 that, the intrinsic compression lines for the samples of the coastal soils investigated compared favorably with that proposed by Burland (1990). Intrinsic compression line for Dhaka clay (Kamaluddin, 1990) also agreed with that proposed by Burland (1990).



**FIGURE 6 : Intrinsic Compression Lines for Samples of Four Coastal Soils**

**Table 6**  
**Intrinsic Properties of the Soils of Four Locations of Coastal Region**

Soil	$w_L$ (%)	$e_L$	$e^*_{100}$	$C_c^*$	Equation for $I_v$
G1	59	1.65	1.02	0.405	$I_v = (e - 1.02)/0.405$
G3	41	1.11	0.73	0.242	$I_v = (e - 0.73)/0.242$
KA	45	1.19	0.80	0.295	$I_v = (e - 0.80)/0.295$
M	66	1.80	0.73	0.272	$I_v = (e - 0.73)/0.272$

**Table 7**  
**Coefficient of Permeability for Reconstituted Samples for Four Coastal Soils Determined from One-dimensional Test**

Soil Type	Average $\sigma'$ (kPa)	Average water content (%)	Average void ratio	$k$ ( $10^{-10}$ m/sec)	
				log t Fitting Method	Fitting Method
G1	40.2	40	1.120	4.70	10.60
	80.4	37	1.035	3.90	8.90
	160.9	33	0.921	3.10	6.90
	321.8	29	0.800	1.90	4.10
	643.5	24	0.669	1.20	2.70
	1072.5	21	0.591	0.80	2.00
G3	40.2	30	0.804	4.98	7.14
	80.4	28	0.757	4.63	6.33
	160.9	26	0.694	3.48	5.26
	321.8	23	0.624	2.63	3.65
	643.5	20	0.549	1.91	2.74
	1072.5	18	0.485	1.28	1.69
KA	40.2	32	0.840	5.25	8.23
	80.4	31	0.809	4.95	7.70
	160.9	29	0.758	4.54	6.94
	321.8	26	0.684	3.98	5.85
	643.5	22	0.587	3.31	4.62
M	40.2	29	0.804	3.53	5.41
	80.4	28	0.758	2.95	4.68
	160.9	25	0.691	2.25	3.33
	321.8	2	0.610	1.42	2.07
	643.5	19	0.523	0.86	1.18
	1072.5	17	0.455	0.53	0.67

**Table 8**  
**Coefficient of Permeability for Undisturbed Samples of Four Coastal Soils**  
**determined from One-dimensional Consolidation Test**

Soil Type	Average $\sigma'$ (kPa)	Average water content (%)	Average void ratio	k ( $10^{-10}$ m/sec)	
				log t Fitting Method	Fitting Method
G1	40.2	36	0.997	18.00	60.00
	80.4	35	0.965	13.00	44.12
	160.9	33	0.908	11.52	33.10
	321.8	30	0.824	6.83	9.50
	643.5	26	0.715	2.78	4.20
G3	40.2	48	1.292	12.00	14.30
	80.4	45	1.207	8.40	19.80
	160.9	40	1.097	7.50	12.40
	321.8	36	0.972	5.30	13.00
	643.5	31	0.848	3.80	4.10
KA	40.2	28	0.747	20.00	54.50
	80.4	27	0.726	18.00	40.00
	160.9	26	0.694	18.30	45.80
	321.8	24	0.648	7.30	24.20
	643.5	22	0.592	5.50	7.70
M	40.2	27	0.747	15.30	99.00
	80.4	26	0.718	13.00	34.90
	160.9	25	0.676	9.60	27.00
	321.8	23	0.620	4.20	13.90
	643.5	20	0.557	3.00	10.80

### *Permeability Characteristics*

Summary of coefficient of permeability derived from consolidation test results for reconstituted samples is presented in Table 7. Permeability determined using log t method and  $\sqrt{t}$  method varies from  $0.53 \times 10^{-10}$  m/s to  $5.25 \times 10^{-10}$  m/s and  $0.67 \times 10^{-10}$  m/s to  $10.60 \times 10^{-10}$  m/s respectively. The values of coefficient of permeability determined using  $\sqrt{t}$  fitting method are higher than those determined using log t fitting method for all the stress ranges. Similar results have also been found by Siddique and Safiullah (1995) for reconstituted Dhaka clay (LL = 40, PI = 20). Table 8 presents summary of coefficient of permeability derived from consolidation test results for undisturbed samples. It can be seen from Table 8 that similar to the reconstituted samples, the values of coefficient of permeability determined using  $\sqrt{t}$  fitting method are higher than those determined using log t fitting method for all the stress ranges.



Permeability determined using  $\log t$  method and  $\sqrt{t}$  method varied from  $3 \times 10^{-10}$  m/s to  $20 \times 10^{-10}$  m/s and  $4 \times 10^{-10}$  m/s to  $99 \times 10^{-10}$  m/s respectively. Comparing the values of permeability of reconstituted and undisturbed samples, it can be seen from Tables 7 and 8 that, the values of permeability for the undisturbed samples of all the coastal soils are significantly higher than those for the reconstituted samples.

Figures 7 and 8 show the plottings of void ratio against logarithm of permeability determined from consolidation tests for reconstituted and undisturbed samples respectively. It can be seen from Figs. 7 and 8 that permeability decreased with decreasing void ratio and that  $e - \log k$  relationships are non-linear. Non-linear  $e - \log k$  relations have been reported by Samarasinghe et al. (1982) and Raymond (1966) for reconstituted normally consolidated clays. Non-linear  $e - \log k$  relations have also been reported by Ahmed (1977), Tavenas et al. (1983) and Siddique and Safiullah (1995) for undisturbed natural soft Ranjit clay of Bangkok, Matagami clay of Canada and reconstituted Dhaka clay respectively.

In an attempt to examine the applicability of the theoretical relation (Eqn. 7) as proposed by Samarasinghe et al. (1982),  $\log[k(1+e)]$  versus  $\log e$  have been plotted for the reconstituted samples of four coastal soils. These plottings are shown in Fig. 9. It can be seen from Fig. 9 that approximately linear  $\log[k(1+e)] - \log e$  relationships have been obtained for the coastal soils. Linear  $\log[k(1+e)] - \log e$  relationships have also been obtained for the undisturbed samples of the four coastal soils. This suggests that the permeability prediction model proposed by Samarasinghe et al. (1982) may be applicable for these coastal soils. The permeability parameters  $n$  and  $C$  of the reconstituted and undisturbed samples of the four coastal soils studied were determined from the linear plots of  $\log[k(1+e)]$  versus  $\log e$ . The magnitudes of these parameters together with corresponding correlation coefficients are presented in Tables 9 and 10 respectively for reconstituted and undisturbed samples. It can be seen from Table 9 that for reconstituted samples, permeability parameter  $n$  varied between 1.7 and 4.1, while permeability parameter  $C$  varied from  $7.2 \times 10^{-10}$  m/s to  $27.6 \times 10^{-10}$  m/s. Table 10 shows the values of  $n$  and  $C$  parameters for the undisturbed samples which varied from 3.1 to 9.1, and from  $11.6 \times 10^{-10}$  m/s to  $1264.7 \times 10^{-10}$  m/s respectively. From the data presented in Tables 9 and 10, it is evident that the values of permeability parameters  $n$  and  $C$  for the undisturbed samples of the coastal soils are significantly higher than those for the respective reconstituted samples.

## Conclusions

Compressibility and permeability characteristics of reconstituted and undisturbed samples of coastal soils of Bangladesh have been studied by

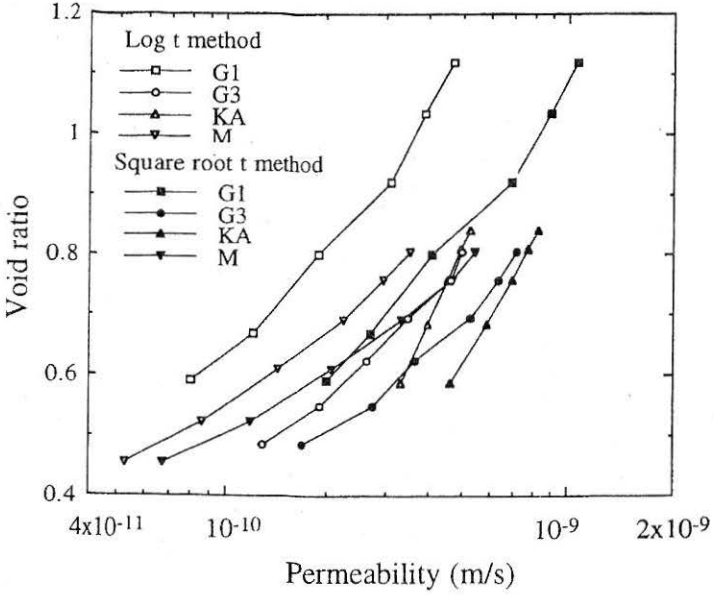


FIGURE 7 : Void Ratio Vs. Log Permeability Plots for Reconstituted Samples of Four Coastal Sites

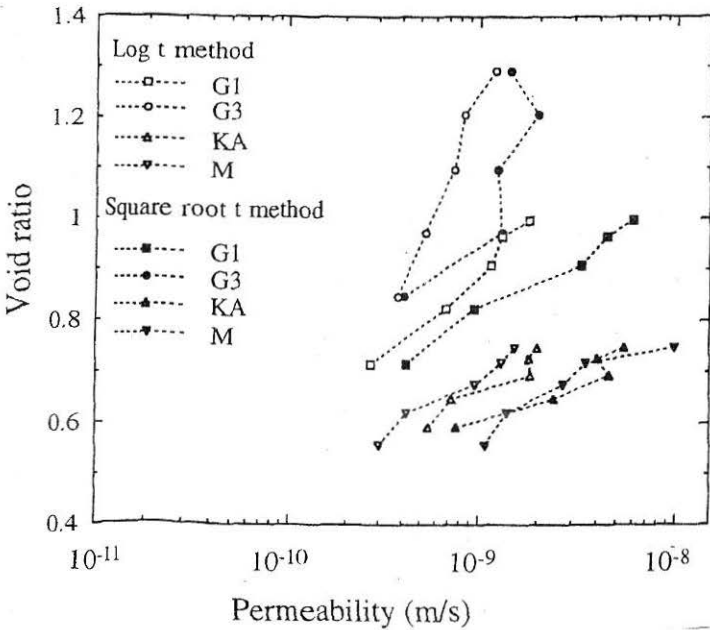


FIGURE 8 : Void Ratio Vs. Log Permeability Plots for Undisturbed Samples of Four Coastal Sites

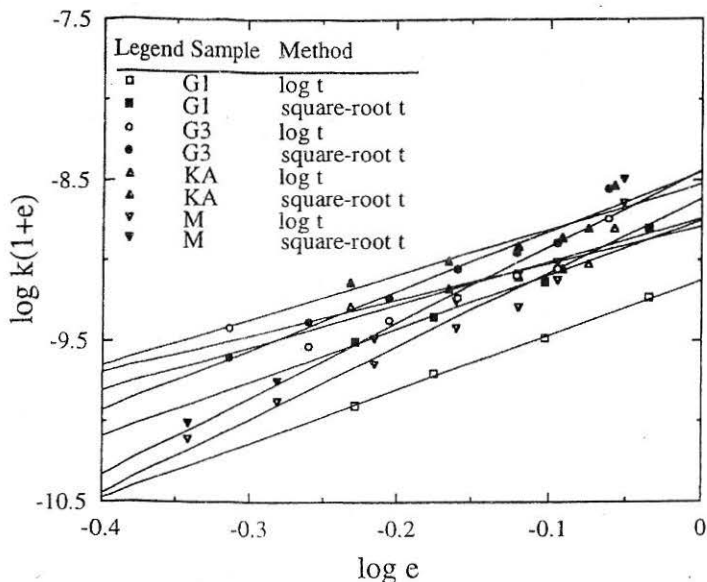


FIGURE 9 :  $\log k(1+e)$  Vs.  $\log e$  Plots for Reconstituted Samples of Four Coastal Soils

performing incremental loading one-dimensional consolidation tests. The major findings and conclusions are summarized below.

The values of compression Index,  $C_c$  and swelling index,  $C_s$  for reconstituted samples vary between 0.29 and 0.40 and 0.028 to 0.046, and for undisturbed samples between 0.22 to 0.42 and 0.037 to 0.062 respectively. Samples of Gohira which has high natural water content the  $C_c$

Table 9  
Values of Permeability Parameters for Reconstituted Samples of Coastal Soils from One-dimensional Consolidation Test

Soil Type	log t Fitting Method			$\sqrt{t}$ Fitting Method		
	n	C ( $10^{-10}$ m/sec)	Correlation Coefficient r	n	C ( $10^{-10}$ m/sec)	Correlation Coefficient r
G1	3.3	7.2	0.999	3.2	16.8	0.992
G3	2.1	13.0	0.909	3.2	27.6	0.997
KA	1.7	12.6	1.0	2.1	21.9	0.997
M	3.9	16.1	0.998	4.1	25.1	1.0

**Table 9**  
**Values of Permeability Parameters for Undisturbed Samples of Coastal Soils**  
**from One-dimensional Consolidation Test**

Soil Type	log t Fitting Method			$\sqrt{t}$ Fitting Method		
	n	C ( $10^{-10}$ m/sec)	Correlation Coefficient r	n	C ( $10^{-10}$ m/sec)	Correlation Coefficient r
G1	6.1	38.2	0.988	9.1	134.6	0.986
G3	3.1	11.6	0.991	3.5	18.6	0.877
KA	6.5	255.3	0.956	8.5	1264.7	0.958
M	6.4	179.5	0.985	7.2	903.6	0.929

values of the undisturbed samples are higher than those of reconstituted samples. For the samples from Kalapara and Mognama which have comparatively low natural water content, however, the  $C_c$  values of the undisturbed samples are lower than those of reconstituted samples. It is also evident that  $C_s$  of undisturbed samples of Gohira are considerably higher than those of reconstituted samples while the  $C_s$  values of undisturbed samples of Kalapara and Mognama are slightly higher than that of reconstituted samples.

It is evident that the values of  $m_v$  and  $m_s$  of reconstituted samples are comparatively higher than undisturbed samples. It is also evident that for the reconstituted samples  $c_v$  value generally increased with increasing levels of consolidation stress. For both reconstituted and undisturbed samples, the values of  $c_v$  obtained using log t method are less than those obtained using  $\sqrt{t}$  method. The values of  $c_v$  for the four reconstituted and undisturbed samples of the coastal soils varied between  $3.3 \times 10^{-4}$  to  $30.2 \times 10^{-4}$  cm<sup>2</sup>/sec and  $10.1 \times 10^{-4}$  to  $224.4 \times 10^{-4}$  cm<sup>2</sup>/sec respectively. For the undisturbed samples, the values of  $c_v$ , in general, reduced with the increase in consolidation stress. The values of  $c_v$  for the undisturbed samples have been found to be significantly higher than those for reconstituted samples.

The range of void ratio at the liquid limit ( $e_L$ ) for four coastal soils is from 1.10 to 1.80, the range of  $e_{100}^*$  is from 0.73 to 1.02 and the range of  $C_c^*$  is from 0.24 to 0.40. The intrinsic compression lines (ICL) for samples of the four coastal soils are in good agreement with the ICL proposed by Burland (1990). The compressibility characteristics of reconstituted clays of the soils studied can be used as a basic frame of reference for interpreting the corresponding characteristics of natural coastal soils. Burland (1990) observed that the sedimentation compression curves for most natural clays lie well above the corresponding intrinsic compression curves. The clays which have these characteristics will be more sensitive and brittle than the reconstituted material.

For reconstituted samples, coefficient of permeability determined using  $\log t$  method and  $\sqrt{t}$  method varied from  $0.53 \times 10^{-10}$  m/s to  $5.25 \times 10^{-10}$  m/s and  $0.67 \times 10^{-10}$  m/s to  $10.60 \times 10^{-10}$  m/s respectively. For undisturbed samples, coefficient of permeability determined using  $\log t$  method and  $\sqrt{t}$  method varied from  $3 \times 10^{-10}$  m/s to  $20 \times 10^{-10}$  m/s and  $4 \times 10^{-10}$  m/s to  $99 \times 10^{-10}$  m/s respectively. Permeabilities computed from consolidation test using  $\sqrt{t}$  fitting method are higher than those determined using  $\log t$  fitting method. The values of permeability for the undisturbed samples of all the coastal soils are significantly higher than those for the reconstituted samples.

The  $e - \log k$  relationships for samples of the four coastal soils have been found to be non-linear. Non-linear  $e - \log k$  relations have been also reported for intact and reconstituted clays (Tavenas et al., 1983; Samarasinghe et al., 1982). However, approximately linear  $\log[k(1+e)] - \log e$  relationships have been obtained for coastal soils. The relation  $k = Ce^n/(1+e)$  as proposed by Samarasinghe et al. (1982), therefore, can be used to predict permeability - void ratio relation for coastal soils. Further research may be carried out using soil samples collected from various other locations of coastal region of Bangladesh in order to confirm the findings of the present work and also to validate the applicability of the above mentioned permeability - void ratio relation in coastal soils. It has been found that compared with the reconstituted samples, the value of permeability parameters  $n$  and  $C$  of the undisturbed samples are considerably higher. For reconstituted coastal samples investigated, permeability parameter  $n$  varied between 1.7 and 4.1, while permeability parameter  $C$  varied from  $7.2 \times 10^{-10}$  m/s to  $27.6 \times 10^{-10}$  m/s. For undisturbed coastal samples, permeability parameter  $n$  varied between 3.1 and 9.1, while permeability parameter  $C$  varied from  $11.6 \times 10^{-10}$  m/s to  $1264.7 \times 10^{-10}$  m/s.

## Notations

$C$	=	permeability parameter
$c_v$	=	coefficient of consolidation
$C_c$	=	compression index
$C_c^*$	=	intrinsic compression index
$C_s$	=	swelling index
$e$	=	void ratio
$e_L$	=	void ratio at liquid limit
$e_0$	=	in-situ void ratio

$e'_1$	=	void ratio at consolidation pressure, $\sigma'_v = 1$ tsf (95.7 kPa)
$e'_{100}$	=	void ratio at consolidation pressure, $\sigma'_v = 100$ kPa
$e'_{1000}$	=	void ratio at consolidation pressure, $\sigma'_v = 1000$ kPa
$G_s$	=	specific gravity
H	=	average length of drainage path
k	=	coefficient of permeability
LL	=	liquid limit
$m_s$	=	coefficient of volume expansion
$m_v$	=	coefficient of volume compressibility
n	=	permeability parameter
PL	=	plastic limit
PI	=	plasticity index
$t_{50}$	=	time corresponding to 50% consolidation
$t_{90}$	=	time corresponding to 90% consolidation
$I_v$	=	void index
$\sigma'_v$	=	vertical effective consolidation stress
$g_w$	=	unit weight of water
$w_1$	=	water content at consolidation pressure, $\sigma'_v = 1$ tsf (95.7 kPa)
$w_L$	=	water content at liquid limit
$w_n$	=	natural water content

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